

Heat Transfer Near An Ice-Ocean Interface In Supercooled Water

Miles McPhee, McPhee Research
Tim Stanton, Naval Postgraduate School
William Shaw, Naval Postgraduate School

With thanks to:

New Zealand National Institute of Water and Atmosphere Research
National Science Foundation Office of Polar Programs

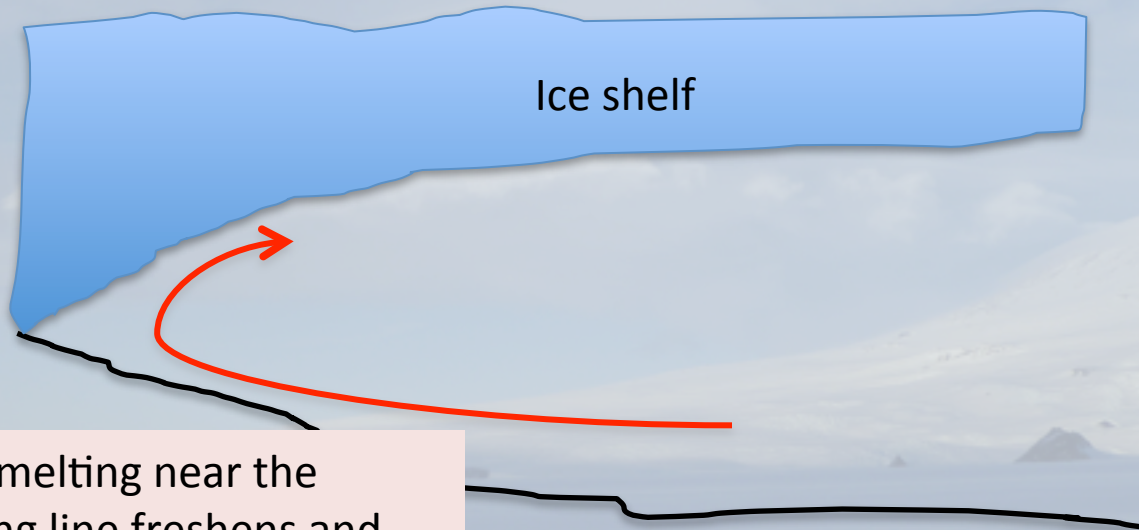
The “ice pump” mechanism



The “ice pump” mechanism

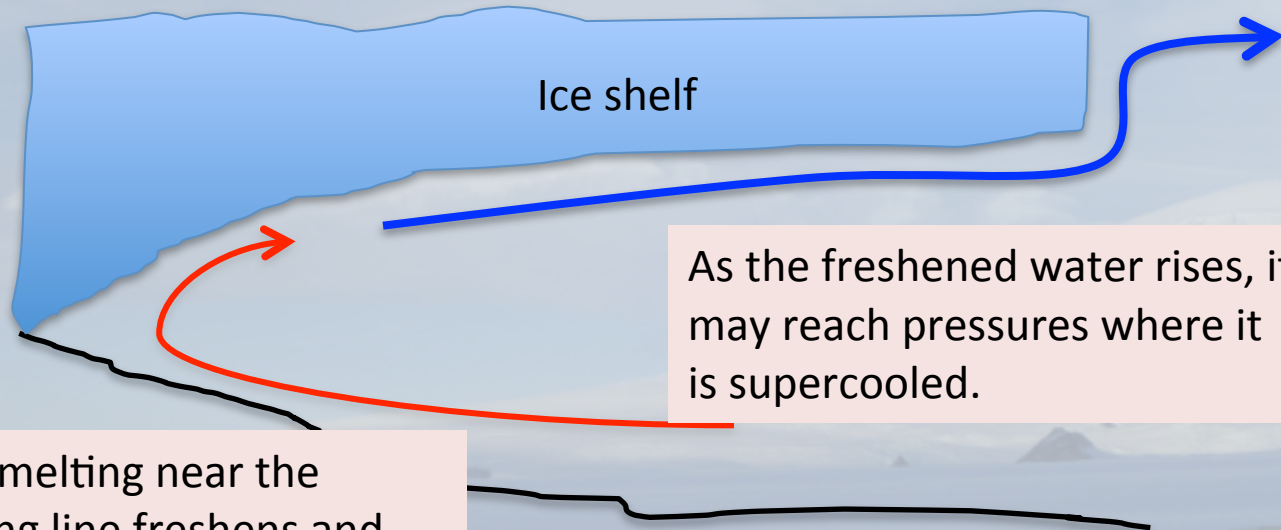


The “ice pump” mechanism



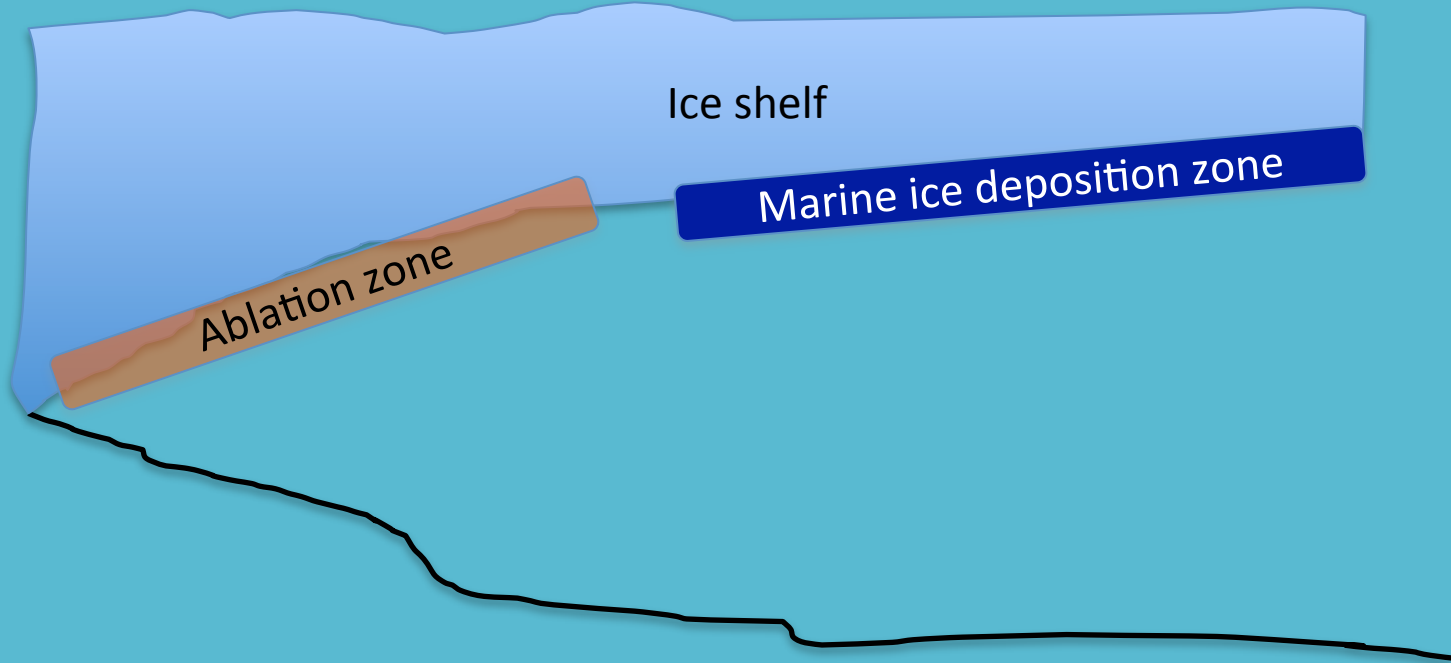
Intense melting near the grounding line freshens and cools incoming deep water.

The “ice pump” mechanism



Intense melting near the grounding line freshens and cools incoming deep water.

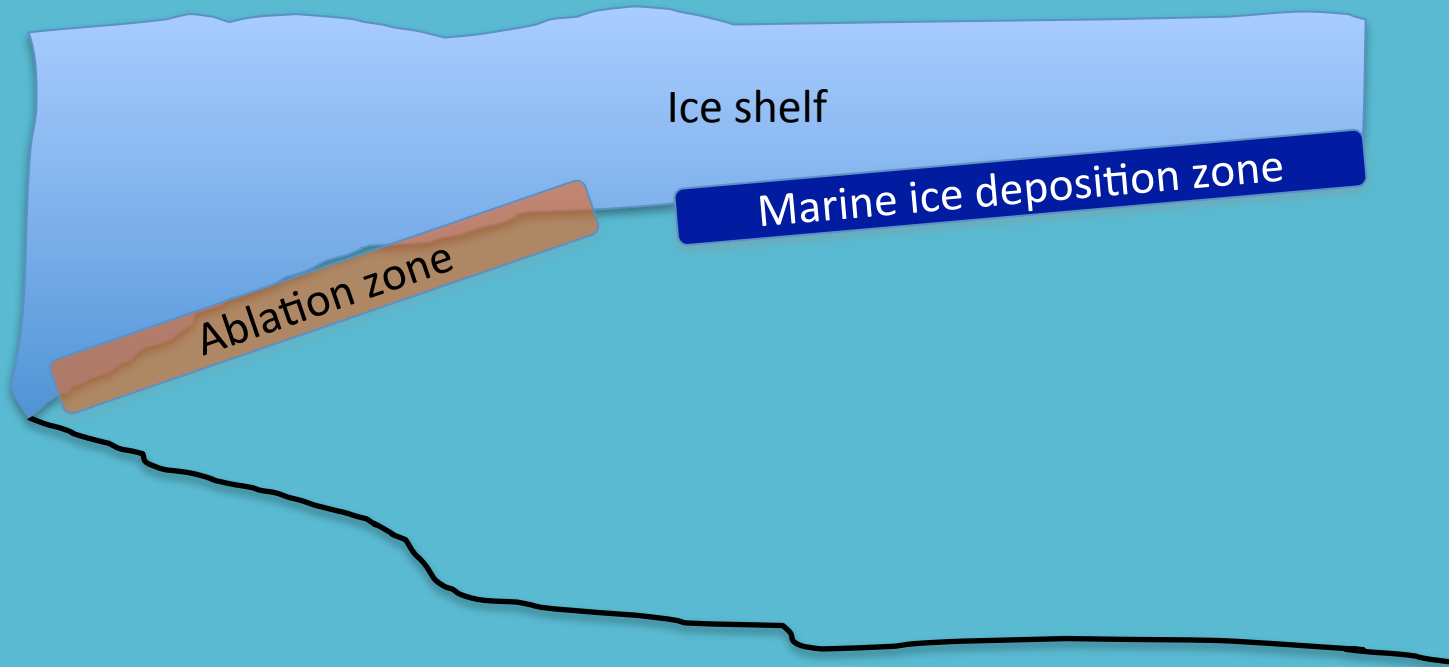
As the freshened water rises, it may reach pressures where it is supercooled.



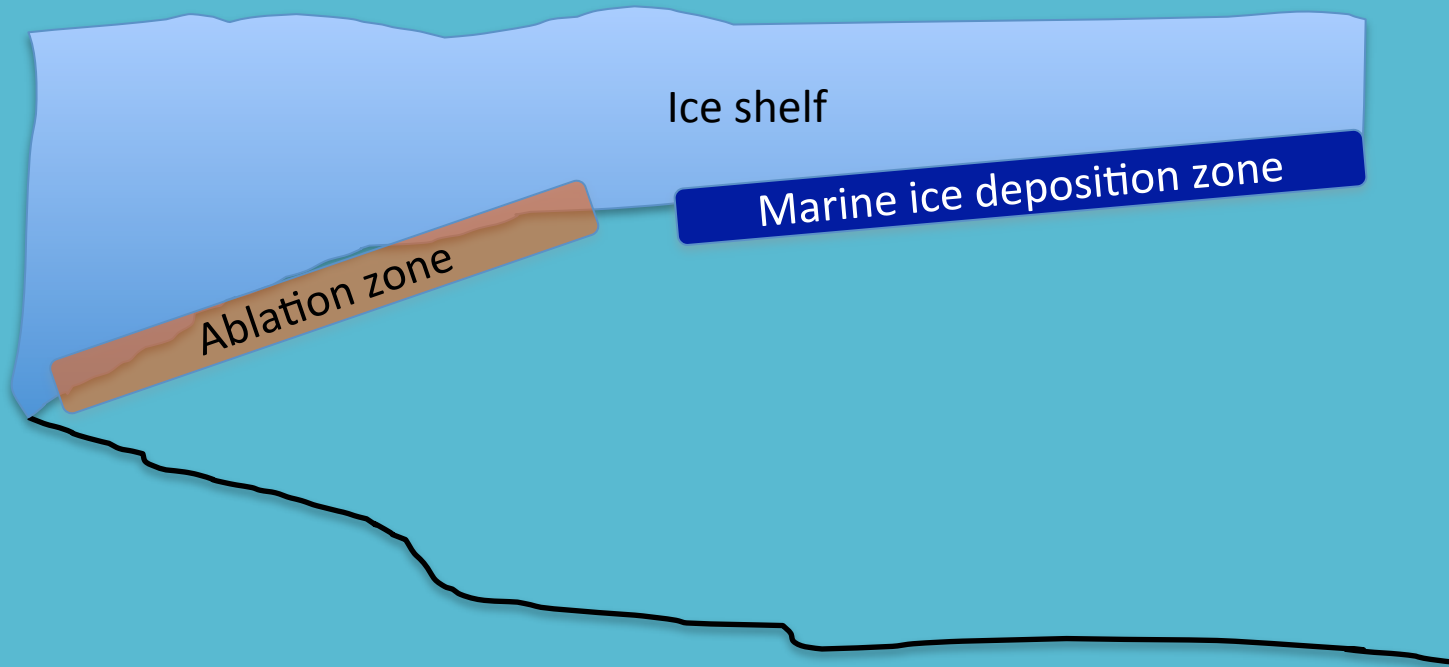
Ice shelf

Ablation zone

Marine ice deposition zone



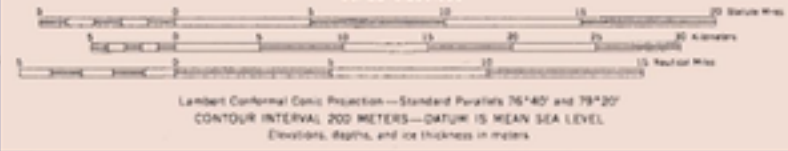
How does basal hydraulic roughness (drag) differ between ablation and deposition zones?



How does basal hydraulic roughness (drag) differ between ablation and deposition zones?

What are the thermodynamics of ice growth in the deposition zone?

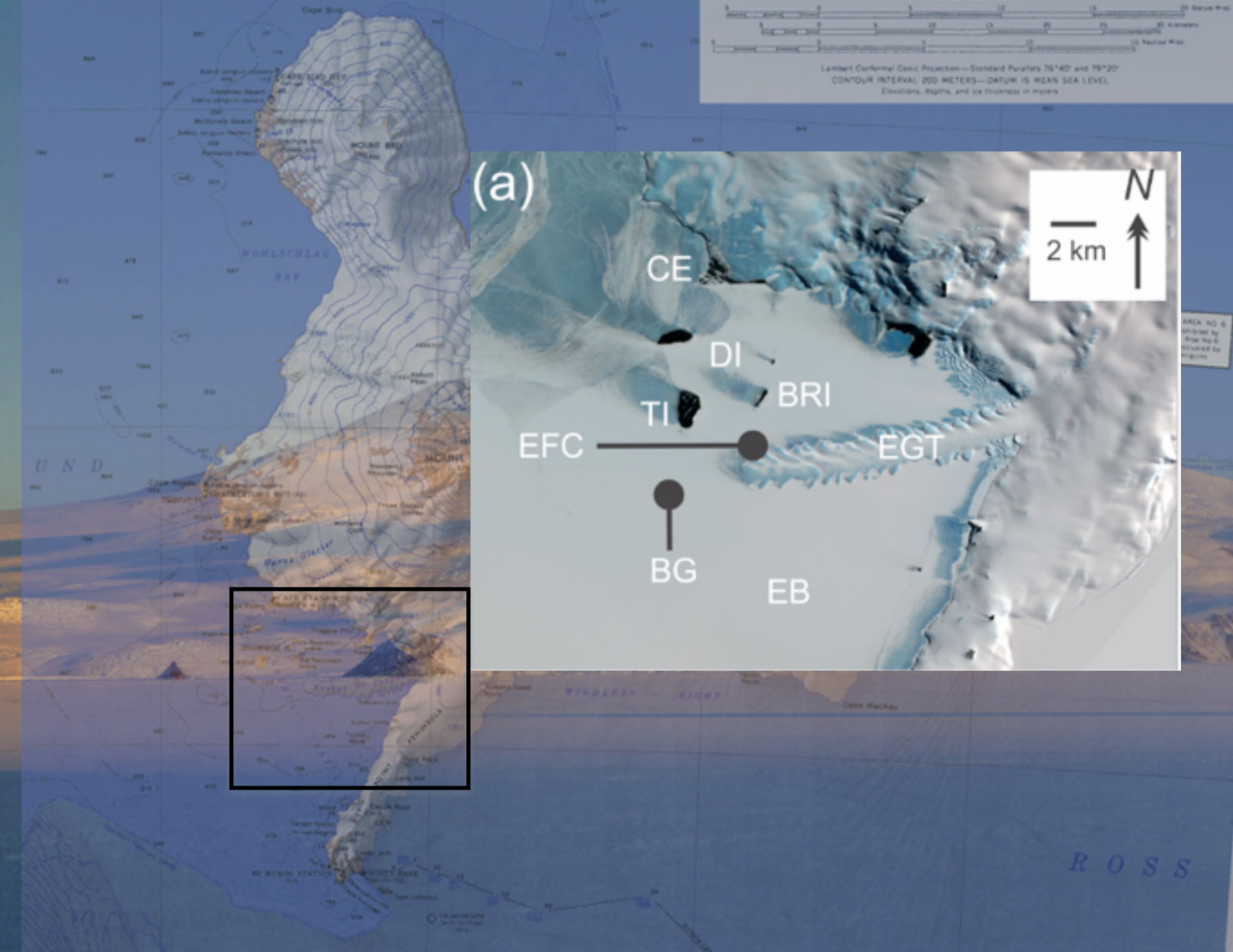




SPECIALLY PROTECTED AREA NO. 6
 Traces or overflight prohibited by international agreement. Area No. 6 also includes any area occupied by the history of Emperor penguins.



ROSS

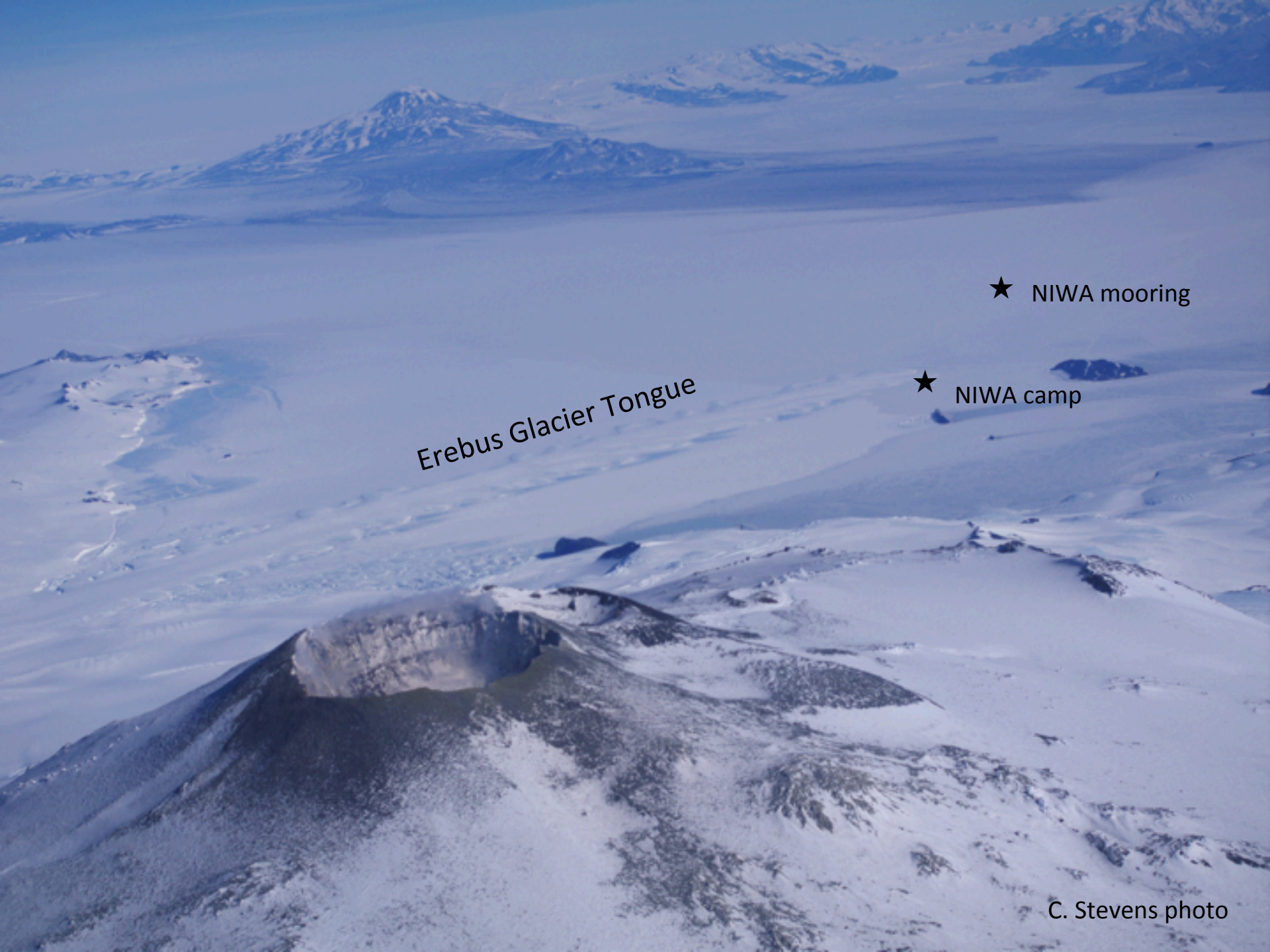


(a)



CE
DI
TI
EFC
EGT
BG
EB

AREA NO. 6
DIVISION BY
AREA NO. 5
COVERED BY
FIGURE 1

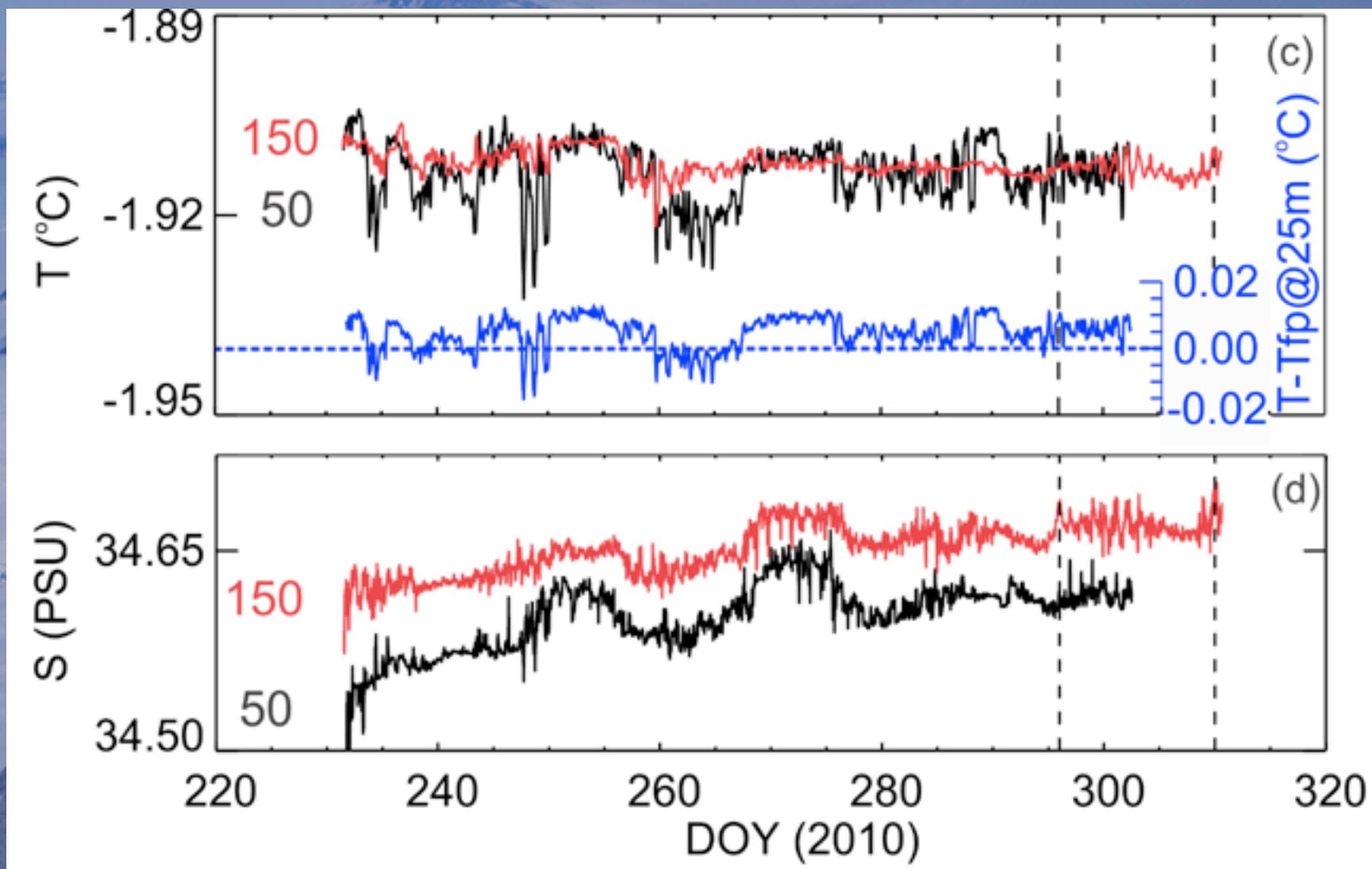


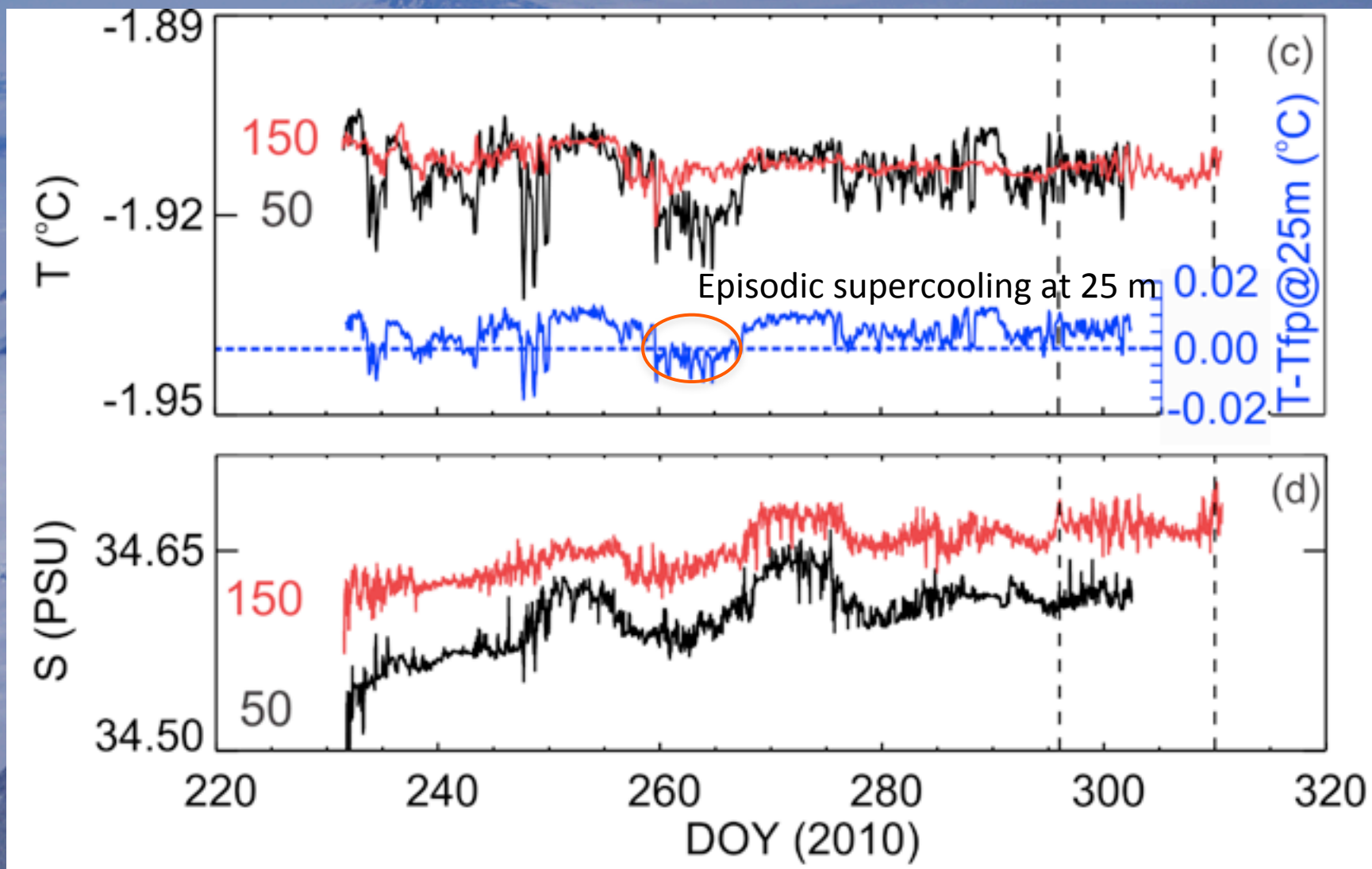
Erebus Glacier Tongue

★ NIWA mooring

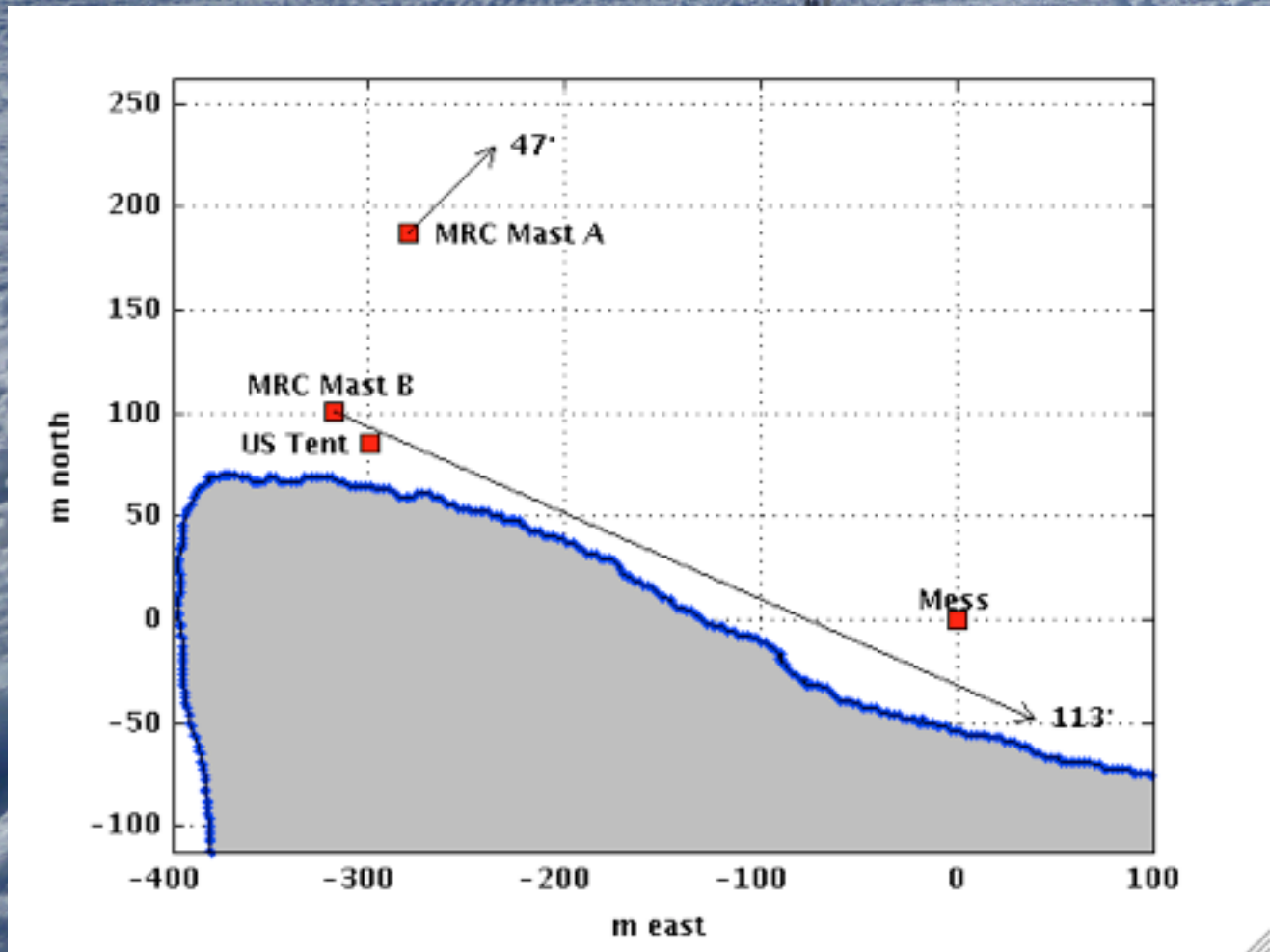
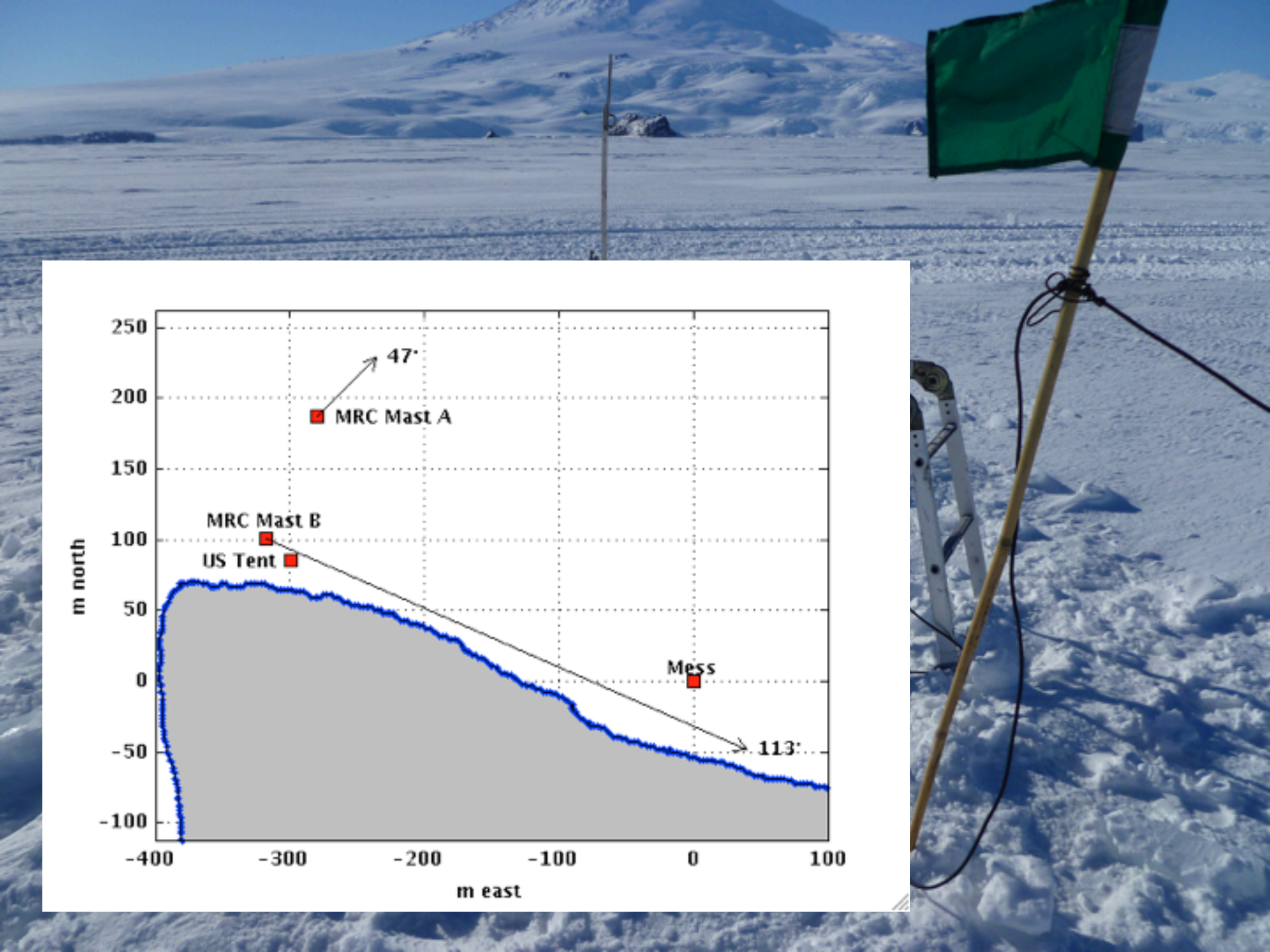
★ NIWA camp

C. Stevens photo



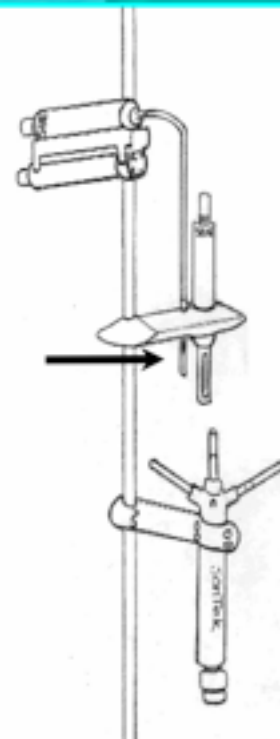








SBE 4 std C

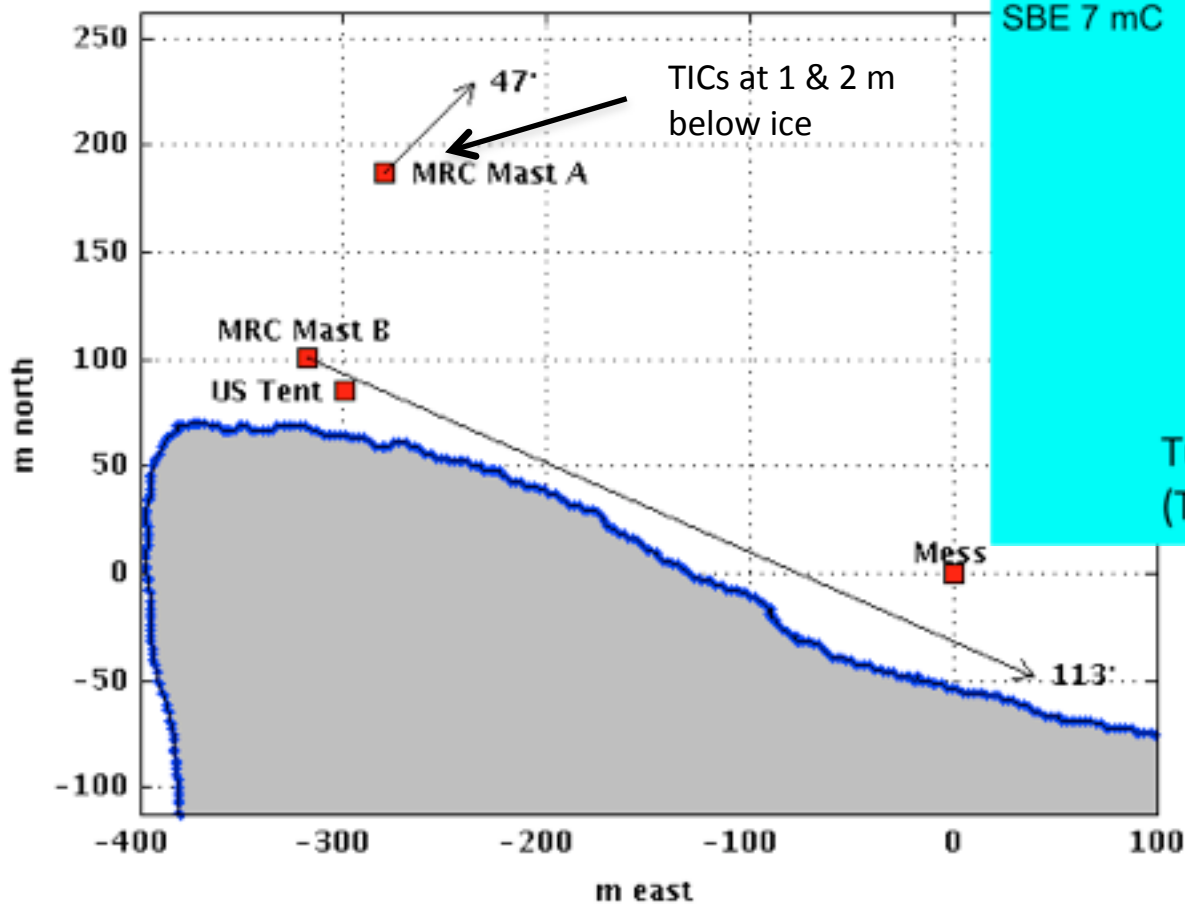


SBE 3 thermometer

SBE 7 mC

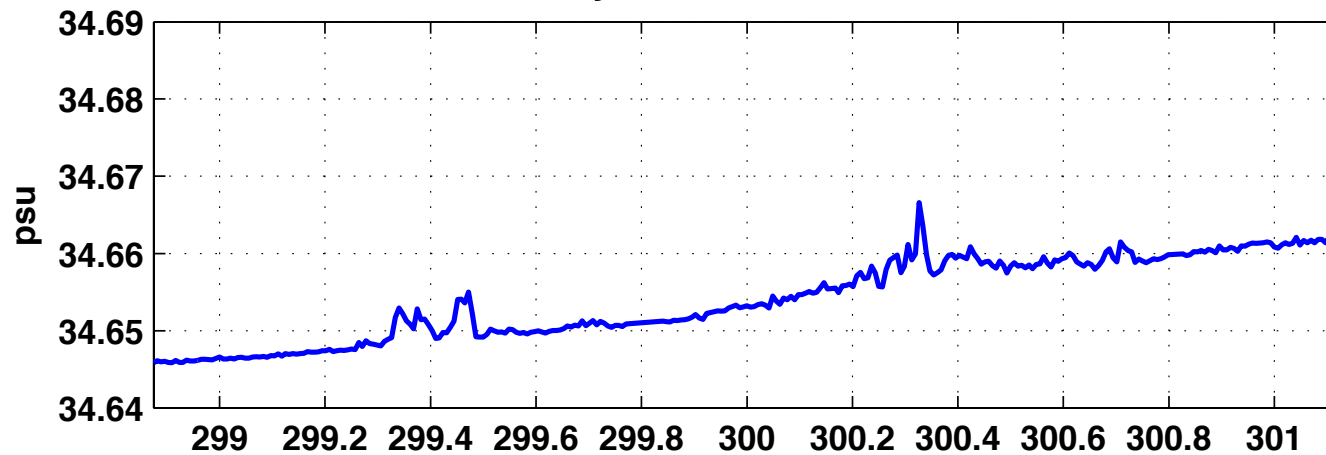
Sontek ADV Ocean (5 Mhz)

Turbulence Instrument Cluster (TIC)

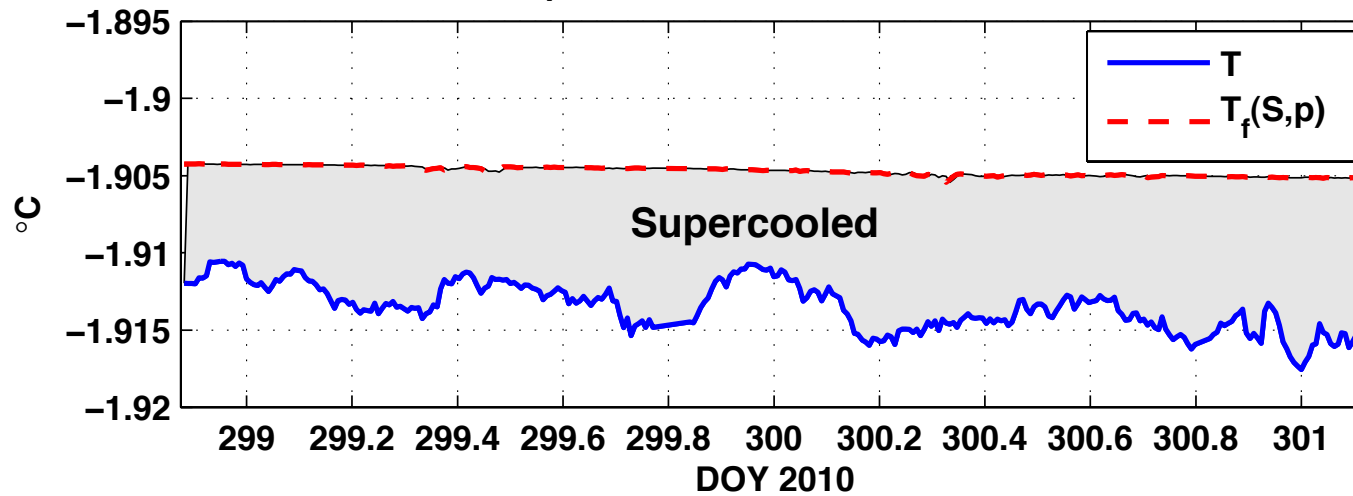




Salinity 1 m below ice at Mast A

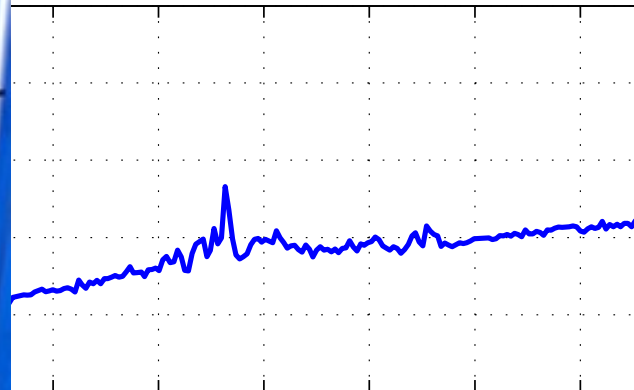


Temperature 1 m below ice at Mast A



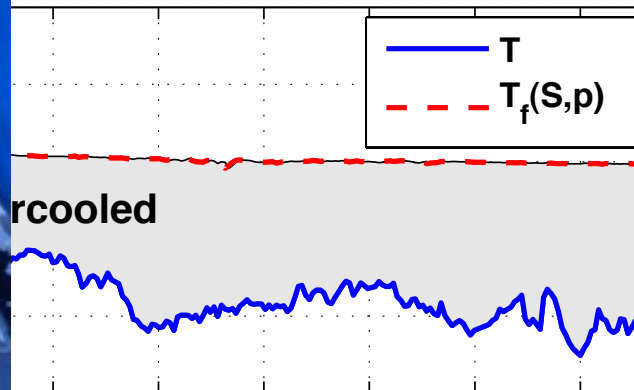


below ice at Mast A



300 300.2 300.4 300.6 300.8 301

n below ice at Mast A



— T
- - - T_f(S,p)

recooled

300 300.2 300.4 300.6 300.8 301

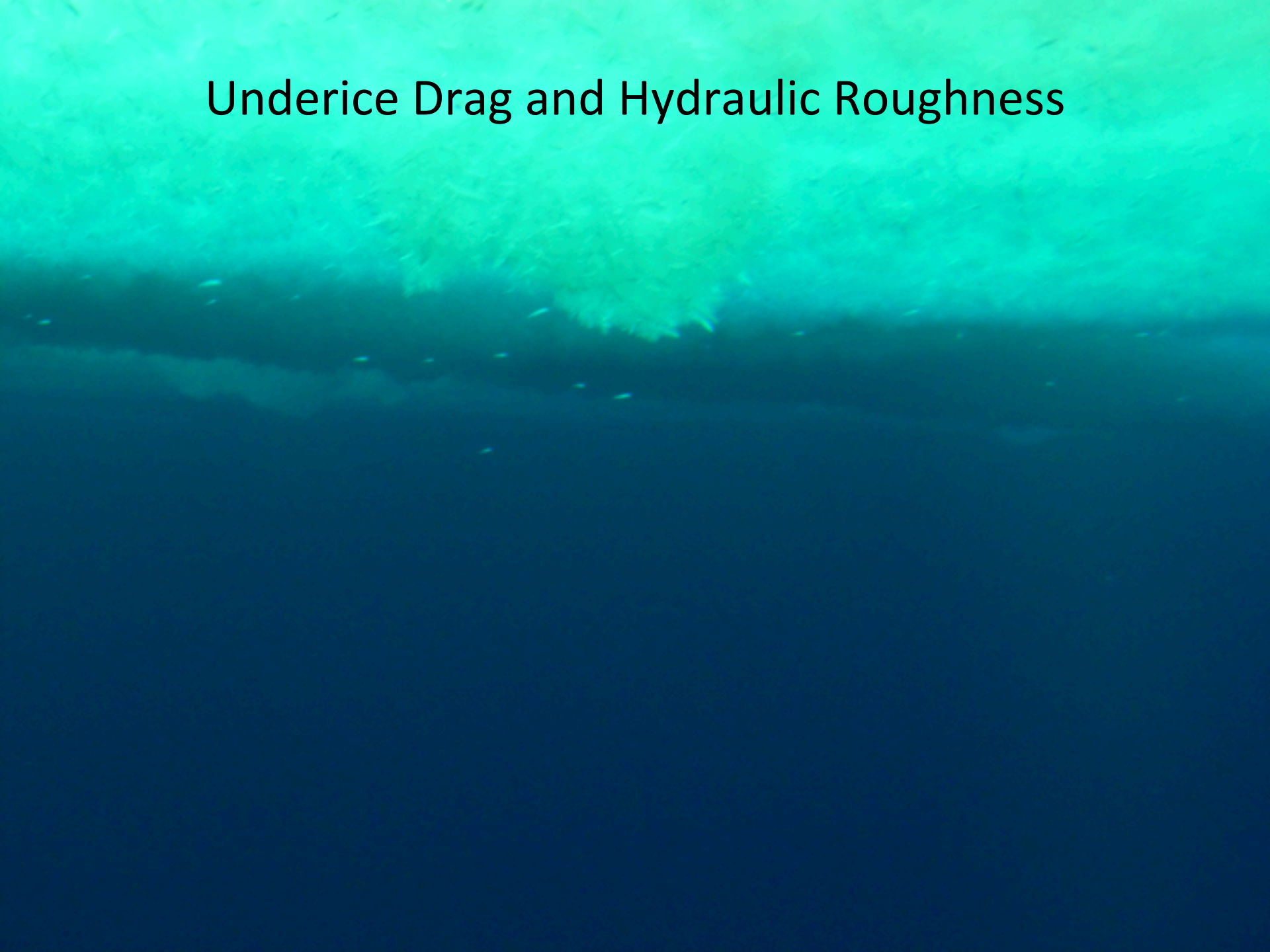
01
DY 2010



below ice at Mast A



Underice Drag and Hydraulic Roughness



Underice Drag and Hydraulic Roughness

Drag Coefficient: kinematic stress proportional to current speed squared

$$\tau = \rho u_*^2 = C_d \rho U^2$$

Underice Drag and Hydraulic Roughness

Drag Coefficient: kinematic stress proportional to current speed squared

$$\tau = u_*^2 = C_d U^2$$

"Law of the wall": $U = (u_* / \kappa) \log(d / z_0) = C_d^{-1/2}$

Measurements 1 m from interface: $\log(z_0) = -\kappa U / u_*$

Underice Drag and Hydraulic Roughness

Drag Coefficient: kinematic stress proportional to current speed squared

$$\tau = u_*^2 = C_d U^2$$

"Law of the wall": $U = (u_* / \kappa) \log(d / z_0) = C_d^{-1/2}$

Measurements 1 m from interface: $\log(z_0) = -\kappa U / u_*$

Previous observations of z_0 under undeformed fast ice:

Barrow Strait (NWT, 1995): 0.03 – 0.05 mm

Vanmijen Fjord, Svalbard, 2001: hydraulically smooth (~0.01 mm)

Freemansundet, Svalbard, 2007: hydraulically smooth (~0.01 mm)



u_* (square root of turbulent stress) from velocity covariance

$$u_{*(\text{cov})} = \left(\langle u'w' \rangle^2 + \langle v'w' \rangle^2 \right)^{1/4}$$

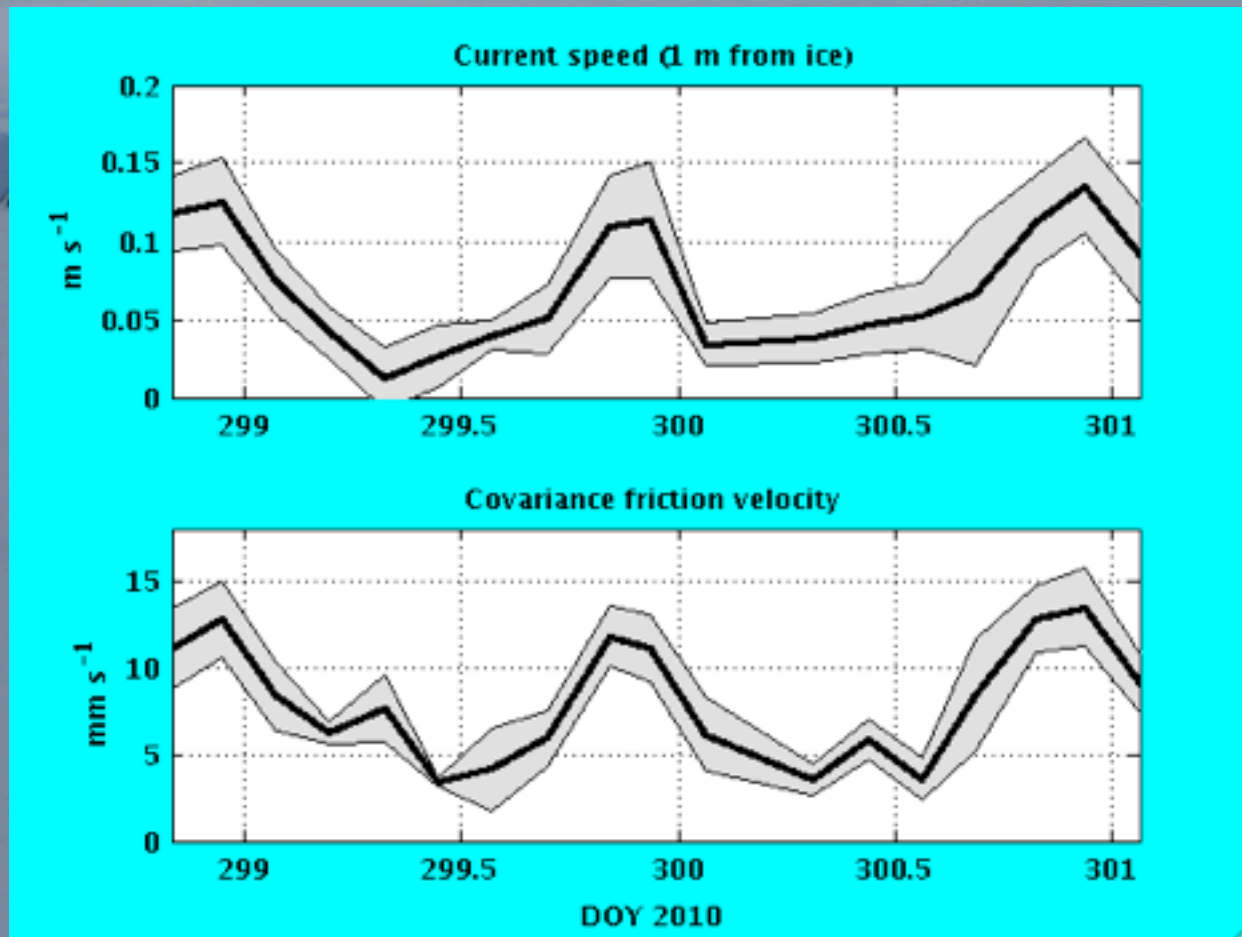
3-hr average of 15 minute turbulent realizations



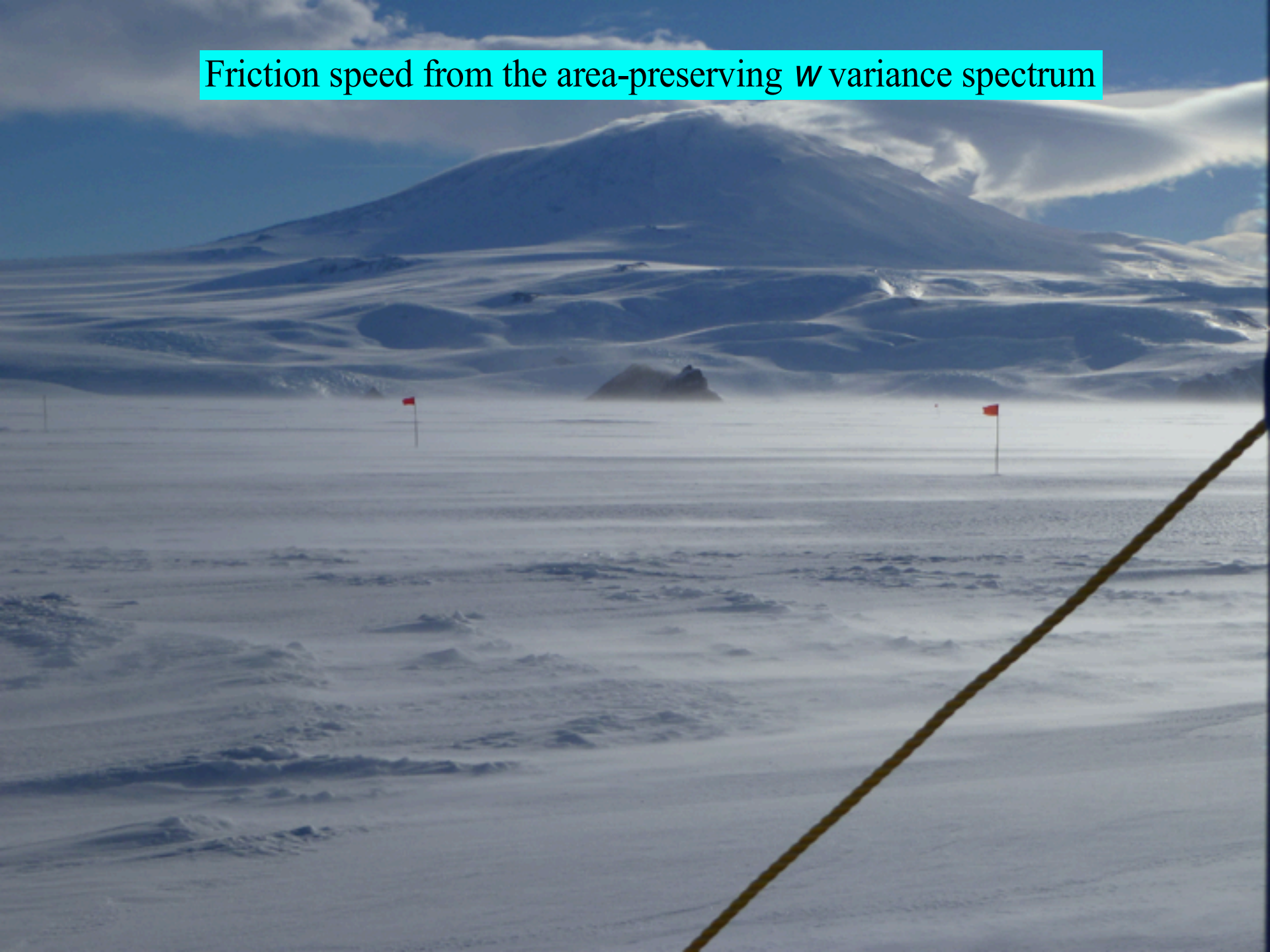
u_* (square root of turbulent stress) from velocity covariance

$$u_{*(\text{cov})} = \left(\langle u'w' \rangle^2 + \langle v'w' \rangle^2 \right)^{1/4}$$

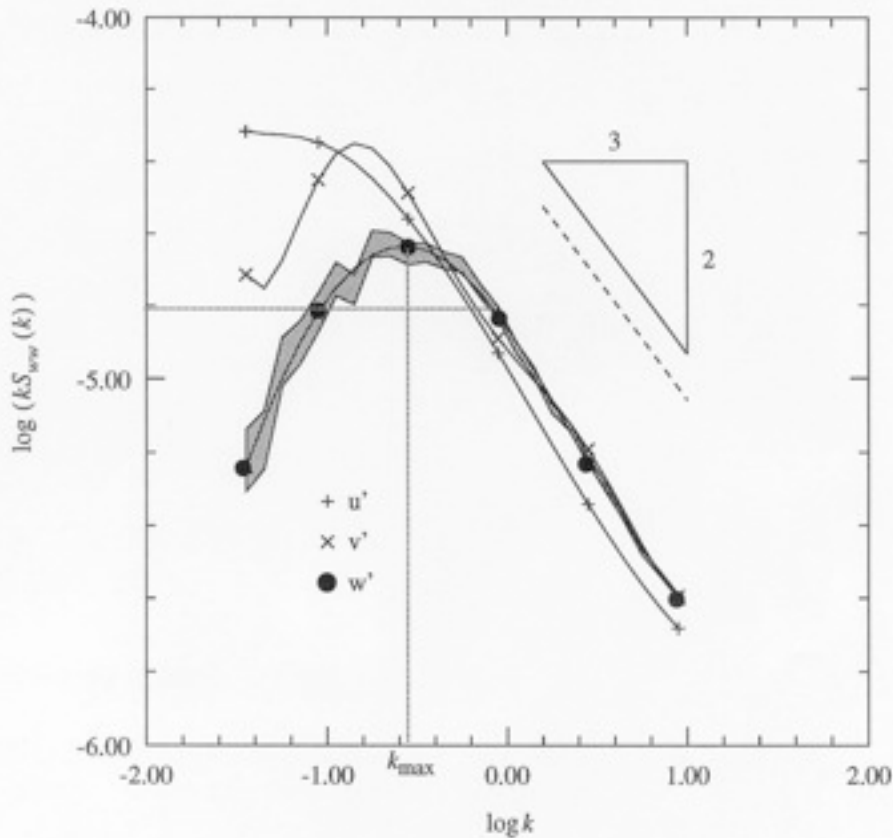
3-hr average of 15 minute turbulent realizations



Friction speed from the area-preserving W variance spectrum



Friction speed from the area-preserving W variance spectrum

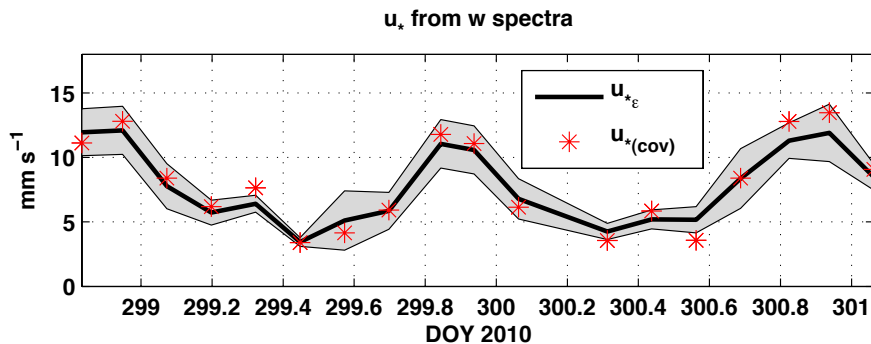
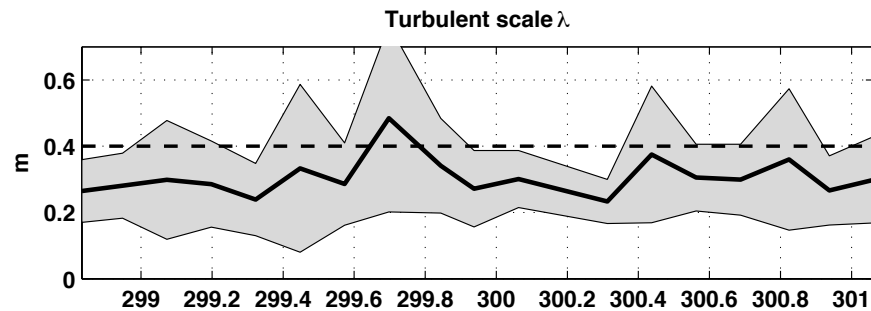
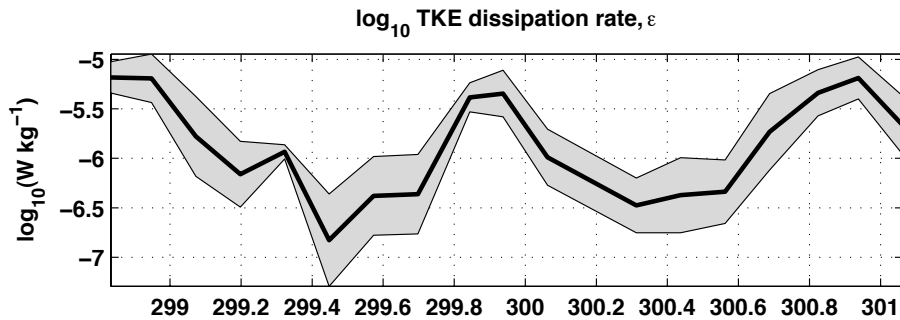


$$kS_{ww}(k) = \frac{4}{3} \alpha_\varepsilon \varepsilon^{2/3} k^{-2/3}$$

$$\lambda = c_\lambda / k_{\max}$$

$$u_{*}(\varepsilon) = (\varepsilon \lambda)^{1/3}$$

Friction speed from the area-preserving w variance spectrum



$$kS_{ww}(k) = \frac{4}{3} \alpha_\epsilon \epsilon^{2/3} k^{-2/3}$$

$$\lambda = c_\lambda / k_{\max}$$

$$u_{*(\epsilon)} = (\epsilon \lambda)^{1/3}$$

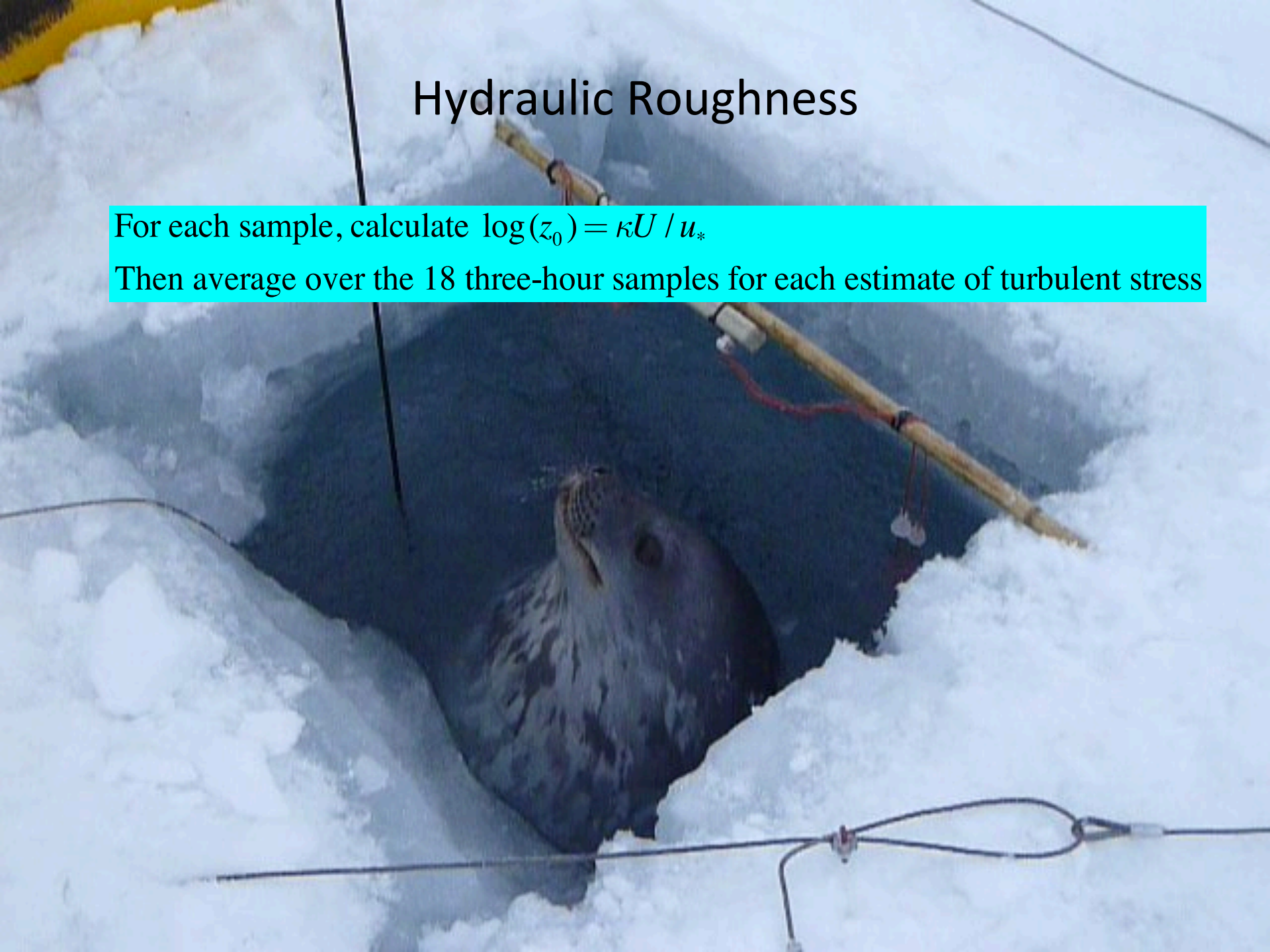
Hydraulic Roughness



Hydraulic Roughness

For each sample, calculate $\log(z_0) = \kappa U / u_*$

Then average over the 18 three-hour samples for each estimate of turbulent stress



Hydraulic Roughness

For each sample, calculate $\log(z_0) = \kappa U / u_*$

Then average over the 18 three-hour samples for each estimate of turbulent stress

From the covariance estimate:

$$\langle \log(z_0) \rangle = -3.51 \pm 1.04 \quad \Rightarrow \quad z_0 = 30 \text{ mm}$$

From the w spectral estimate:

$$\langle \log(z_0) \rangle = -3.47 \pm 0.92 \quad \Rightarrow \quad z_0 = 31 \text{ mm}$$

Hydraulic Roughness

For each sample, calculate $\log(z_0) = \kappa U / u_*$

Then average over the 18 three-hour samples for each estimate of turbulent stress

From the covariance estimate:

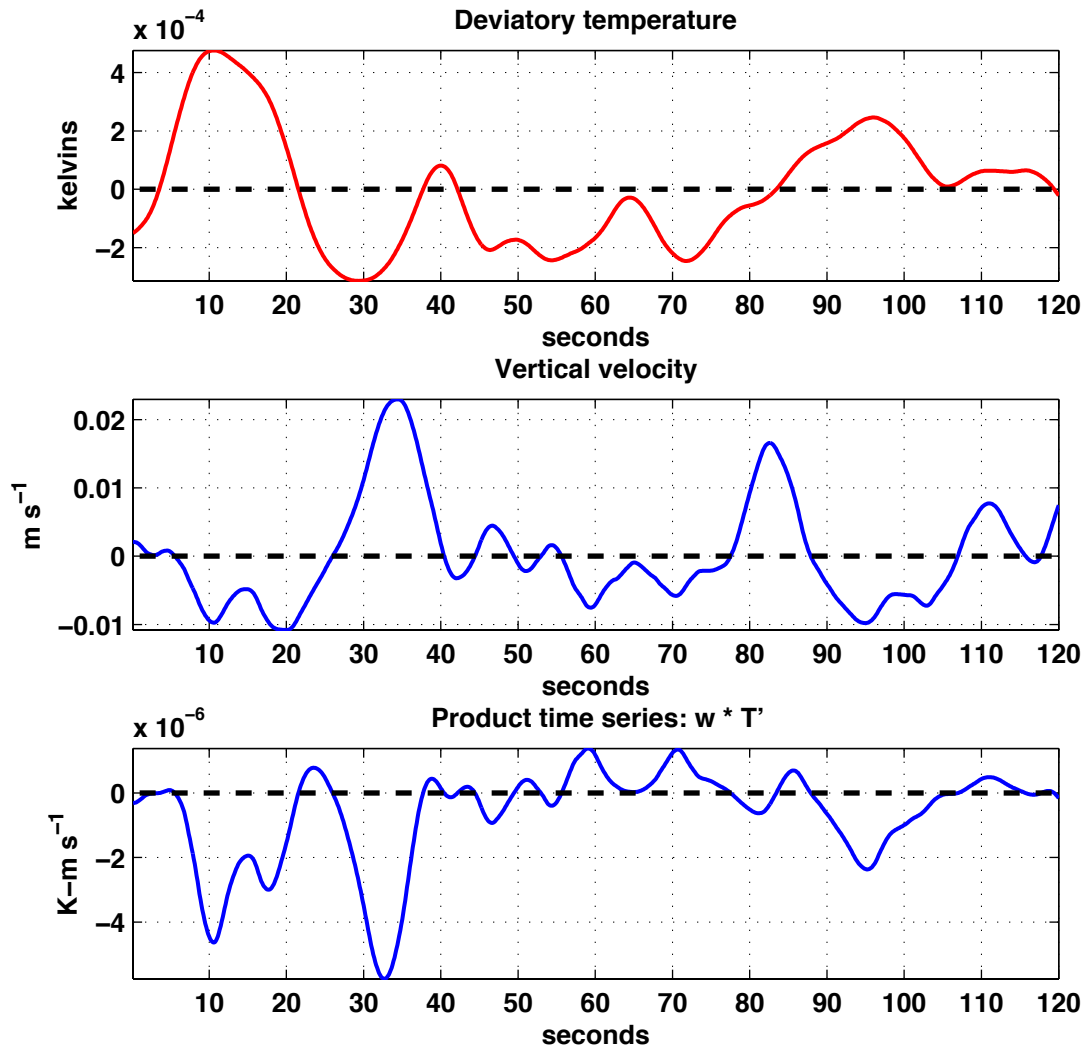
$$\langle \log(z_0) \rangle = -3.51 \pm 1.04 \quad \Rightarrow \quad z_0 = 30 \text{ mm}$$

From the w spectral estimate:

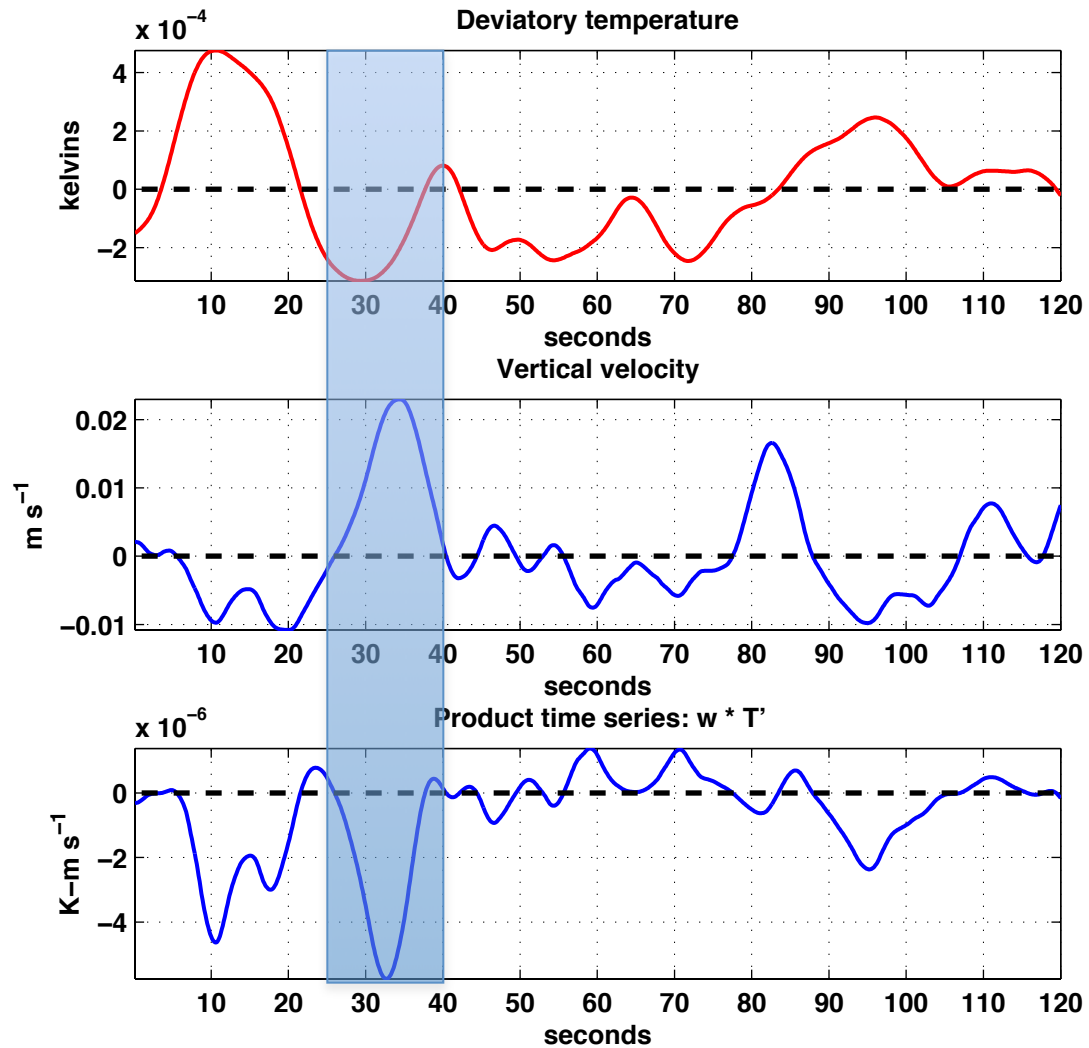
$$\langle \log(z_0) \rangle = -3.47 \pm 0.92 \quad \Rightarrow \quad z_0 = 31 \text{ mm}$$

These values are more typical of deformed, multiyear pack, e.g. Arctic Beaufort, Northpole, western Weddell Sea: 30-60 mm, than for undeformed fast ice without platelet growth (~ 0.01 mm)

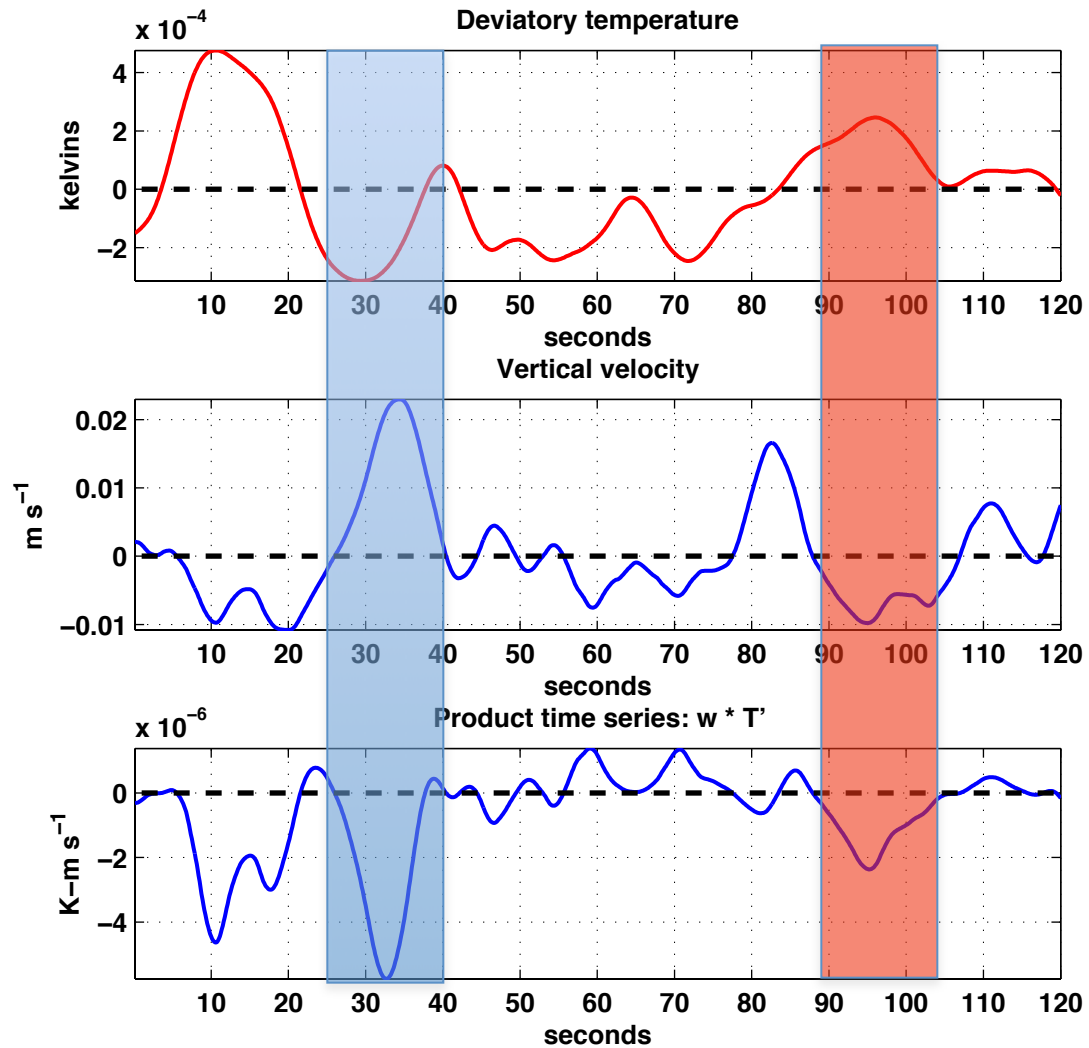
Turbulent Heat Flux

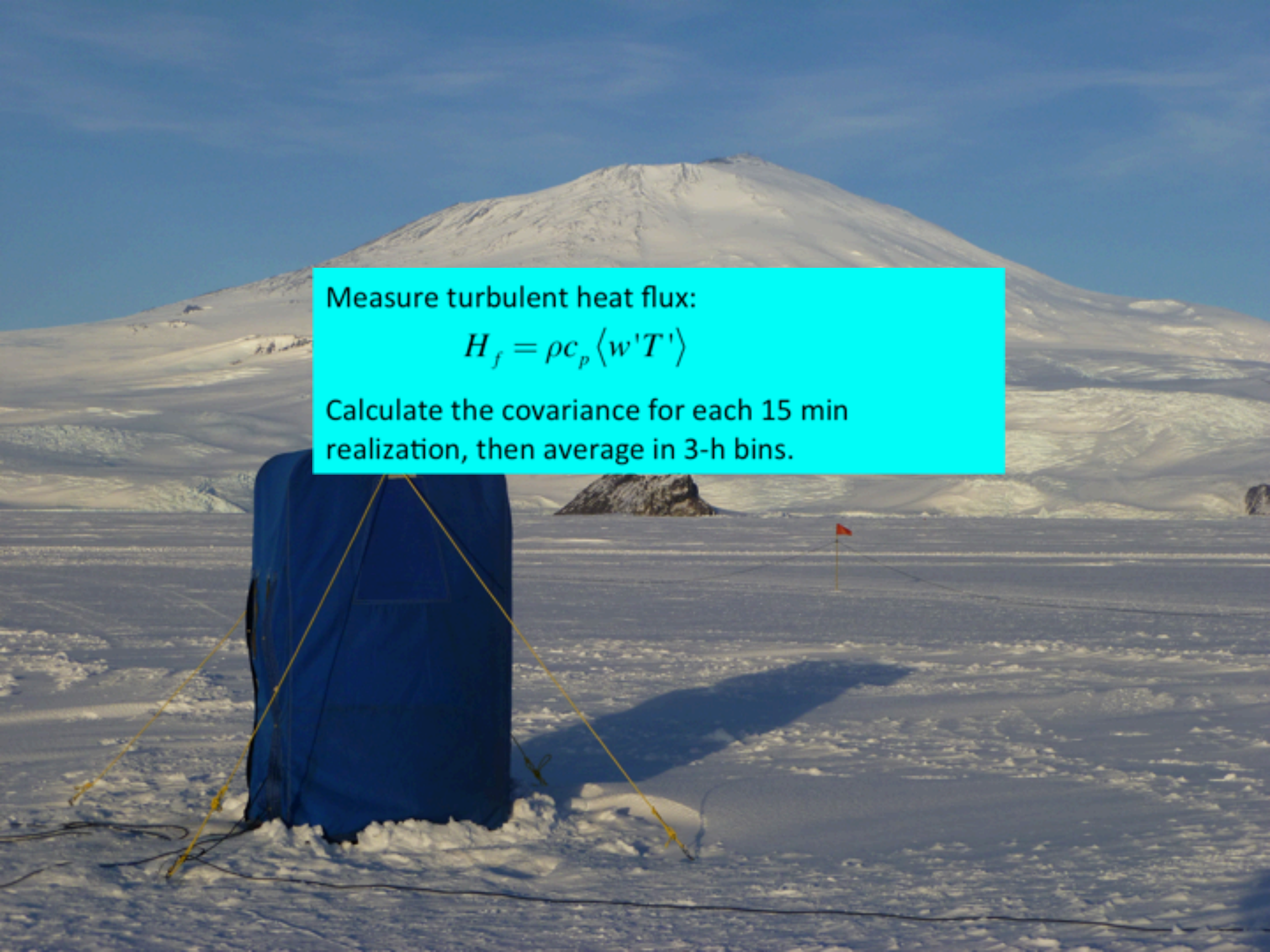


Turbulent Heat Flux



Turbulent Heat Flux

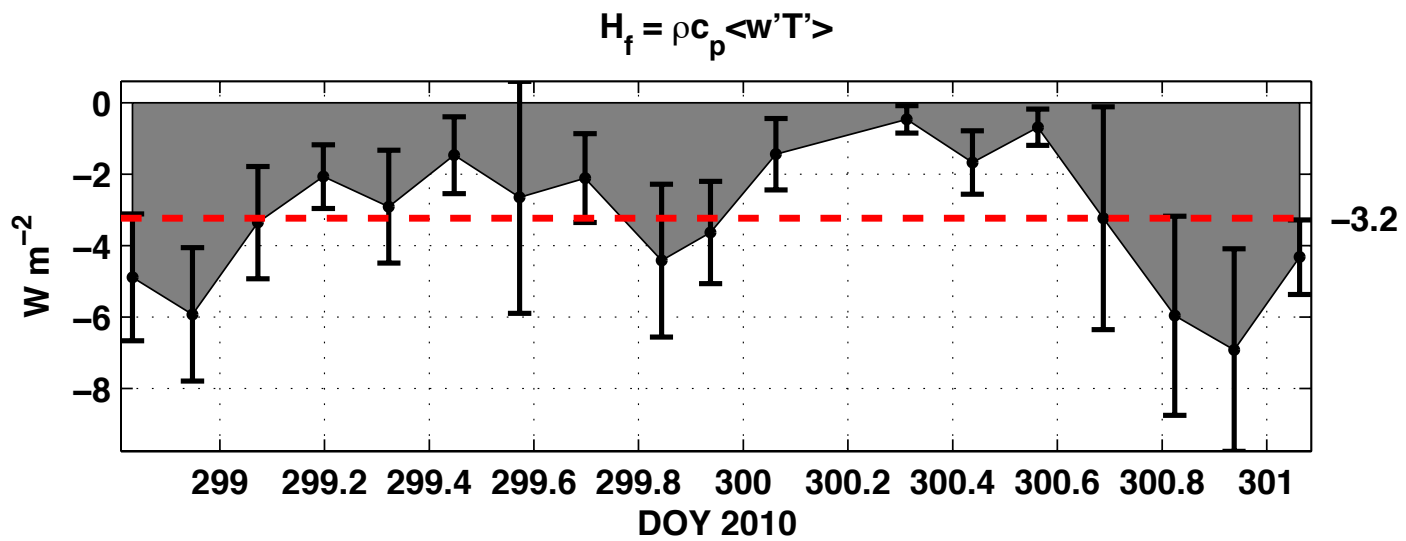
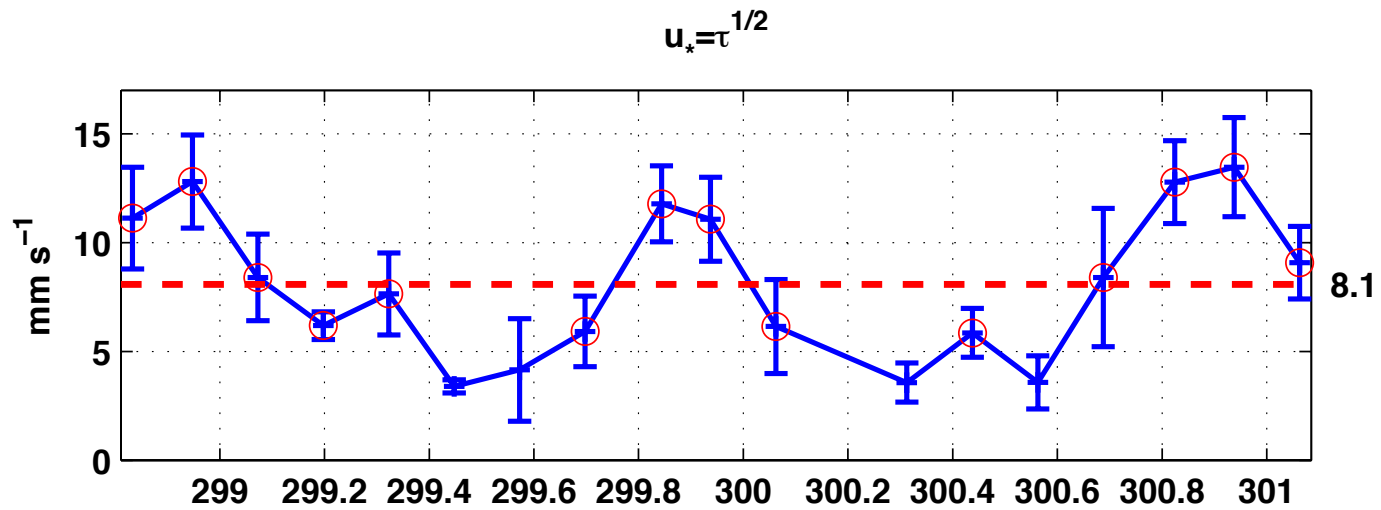




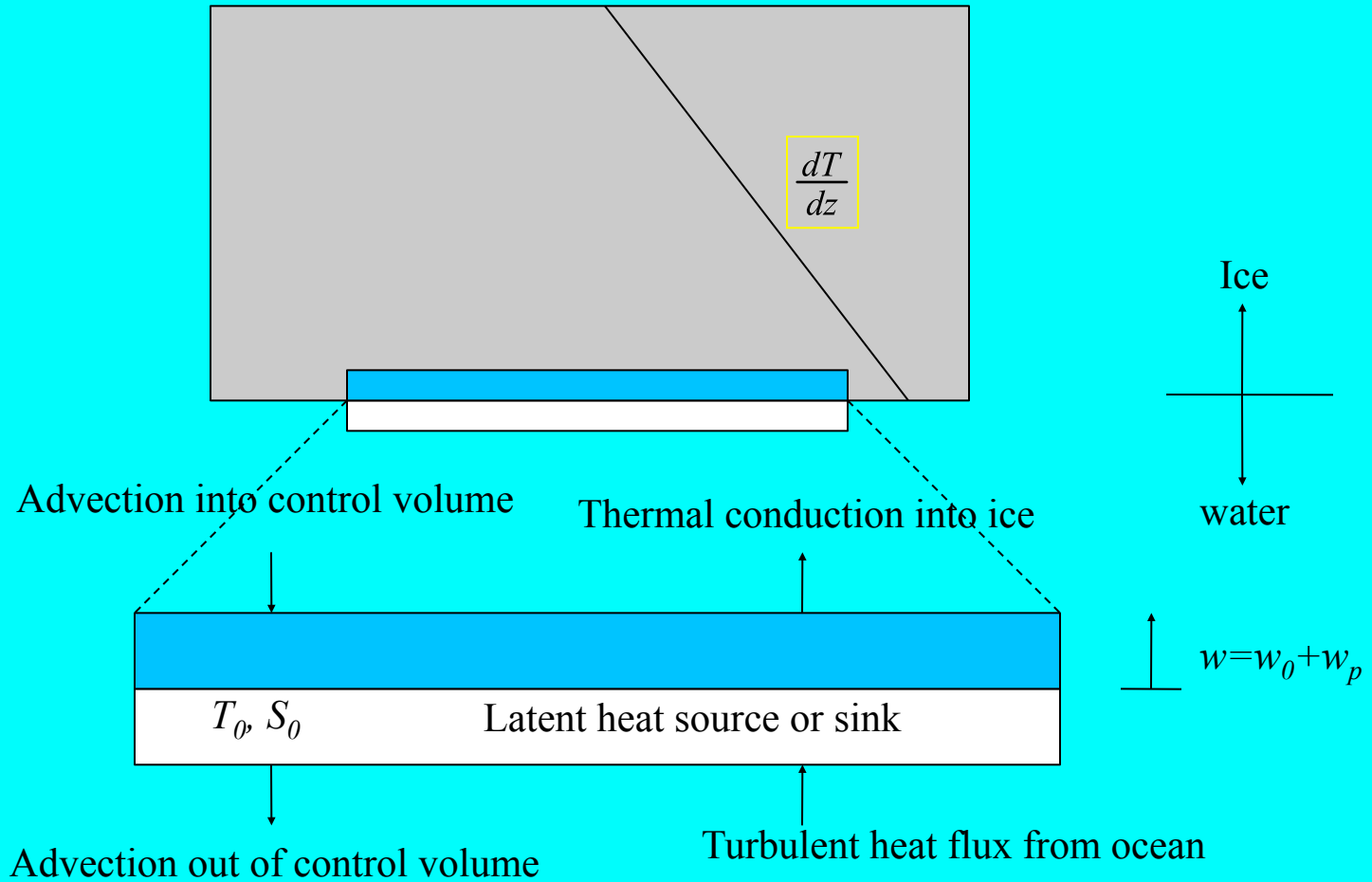
Measure turbulent heat flux:

$$H_f = \rho c_p \langle w'T' \rangle$$

Calculate the covariance for each 15 min realization, then average in 3-h bins.



Thermal Balance at the Ice/Ocean Interface



Dimensional analysis: $\langle w'T' \rangle \propto u_* \delta T$

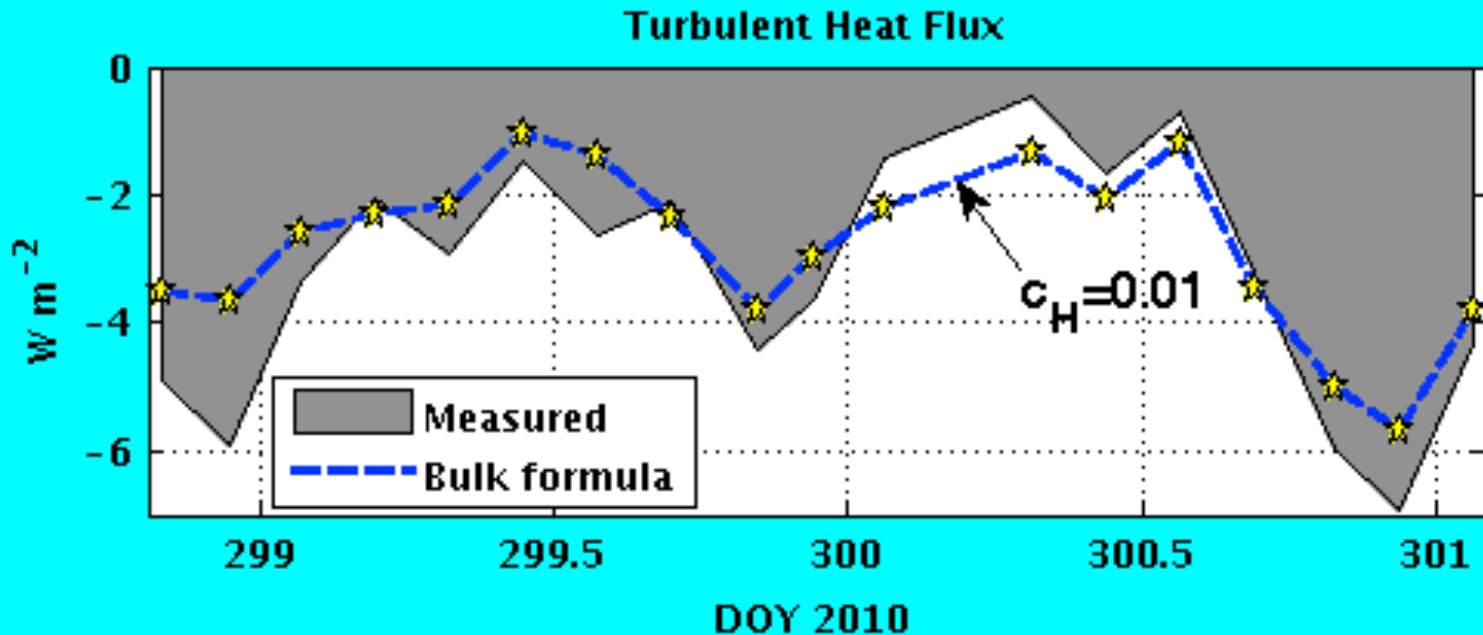
and we find for moderate melting that $H_f \approx \rho c_P c_H u_* [T_{ml} - T_f(S_{ml})]$

where $c_H \sim 0.006$

Dimensional analysis: $\langle w'T' \rangle \propto u_* \delta T$

and we find for moderate melting that $H_f \approx \rho c_P c_H u_* [T_{ml} - T_f(S_{ml})]$

where $c_H \sim 0.006$



Summary

- McMurdo Sound is an excellent laboratory for studying ice/ocean interaction in supercooled water
- The persistence of supercooling in the sound suggests that the water column lacks sufficient nucleation sites for extensive frazil production
- Platelet growth appears to increase the hydraulic roughness (drag) of the ice underside compared with more typical undeformed fast ice
- Platelet growth on the ice underside appears to be rate limited by turbulent heat transport away from the boundary, which depends on both u_* and ΔT

